

PhyTTA: Physics-Informed Test-Time Adaptation of Foundation Models for Regional Drought Prediction

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Abstract

Drought prediction is crucial for disaster mitigation, yet it remains challenging because regional droughts reflect nonlinear interactions among precipitation supply, evaporative demand, and slowly varying land atmosphere states. Although time series foundation models (TSFMs) have shown strong zero shot performance in general forecasting tasks, they can exhibit systematic sensitivity mismatch in the prediction of regional precipitation and temperature anomalies in the Standardized Precipitation Evapotranspiration Index (SPEI). We introduce PhyTTA, a physics informed test time adaptation framework that keeps the TSFM backbone frozen and corrects the predictable part of its one step residual through three components: a dynamic forcing correction, a static tracker of slowly varying background bias, and a conditional confidence gate that applies the correction only in supported regimes. Our analysis shows that, when frozen model residuals contain a component predictable from local forcing changes or slowly varying background bias, correcting this component reduces expected squared error in the population setting. PhyTTA implements this principle through a conservative online ridge EMA correction with regime support gating. Across four foundation backbones on three South Australian locations and TimesFM on three cross regional regions, PhyTTA reduces MSE by up to 19.78%. With the TimesFM backbone it gives consistent 12.3% to 15.1% MSE reductions across the three South Australian locations.

1 Introduction

Drought is one of the most expensive natural disasters, causing billions of dollars in agricultural losses annually and threatening water security for millions of people [Dai, 2011; Trenberth *et al.*, 2014]. The global land area affected by multiyear droughts has expanded steadily over the past four decades [Chen *et al.*, 2025a]. Accurate regional drought prediction is therefore important for water resource manage-

ment, agricultural planning, and early warning systems. The Standardized Precipitation Evapotranspiration Index (SPEI) is widely used for this purpose because it combines precipitation deficits and temperature induced evaporative demand in a single water balance index [Vicente-Serrano *et al.*, 2010; Beguería *et al.*, 2014].

Traditional drought forecasting has relied on statistical models such as Autoregressive Integrated Moving Average (ARIMA) and Seasonal ARIMA (SARIMA), but these models typically assume stationarity and struggle to represent nonlinear climate dynamics [Box *et al.*, 2015; Hyndman and Athanasopoulos, 2021]. Traditional machine learning models reduce some of these restrictions, but they often require extensive feature engineering [Belgiu and Drăguț, 2016]. Deep architectures such as LSTMs [Hochreiter and Schmidhuber, 1997] and Transformers [Vaswani *et al.*, 2017] can learn latent representations directly from data and model long term dependencies, including direct applications to SPEI based drought forecasting [Chen *et al.*, 2025b]. However, training such models from scratch is difficult when high quality regional drought records are limited. This data scarcity constraint motivates the use of pretrained forecasting models rather than purely region specific training.

Recent advances in foundation models appear well suited to this challenge. Time series foundation models (TSFMs) such as TimesFM [Das *et al.*, 2024], Chronos [Ansari *et al.*, 2024], and Moirai [Woo *et al.*, 2024] have demonstrated strong zero shot forecasting across regions. Trained on large heterogeneous corpora, they provide a practical starting point for climate tasks where local data are scarce [Nguyen *et al.*, 2023]. However, when applied to regional drought prediction they can exhibit systematic biases: broad heterogeneous pre training does not necessarily encode localized climate sensitivities driven by region specific soil moisture, vegetation, and evaporative demand processes [Reichstein *et al.*, 2019; Dai, 2011].

The immediate question is therefore not whether to retrain the foundation model, but whether its regional residuals contain a predictable component that can be corrected safely at inference time. Existing time series Test Time Adaptation (TTA) methods such as TAFAS [Kim *et al.*, 2025] and the concurrent PETSA [Medeiros *et al.*, 2025] adapt calibration layers or low rank modules during inference, but

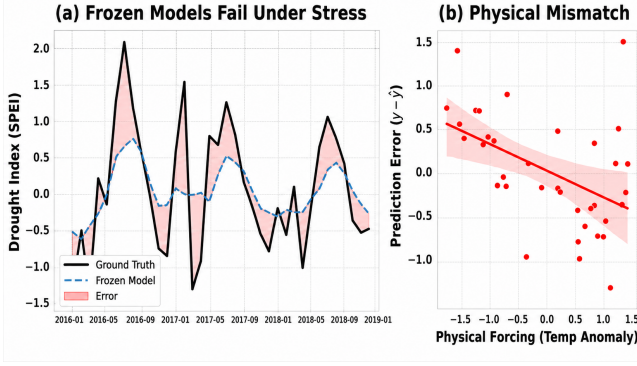


Figure 1: **Motivation for physics informed adaptation.** (a) Frozen foundation models can dampen drought severity fluctuations during extreme events. (b) Prediction residuals can be correlated with physical forcing, here a temperature anomaly, indicating an informed residual component. The figure motivates residual correction conditioned on observable forcing changes, not full hydrological simulation.

gradient updates on small noisy regional streams may degrade performance under distribution shift [Wang *et al.*, 2021; Wilson and Cook, 2020]. We use TAFAS as the representative gradient based benchmark. RevIN [Kim *et al.*, 2022] performs statistical normalization but does not distinguish forcing linked residuals from random weather variability.

In climate science, regional forecasting errors can be addressed better with physical and process understanding, not by statistical alignment alone [Eyring *et al.*, 2024; Reichstein *et al.*, 2019]. Physics informed machine learning work argues that physical constraints and causally informed features can improve robustness and generalization [Karniadakis *et al.*, 2021; Eyring *et al.*, 2024; Runge *et al.*, 2023; Iglesias-Suarez *et al.*, 2024]. We put this perspective into practice by using physical forcing variables, observed climate drivers of the SPEI water balance such as precipitation, maximum temperature, humidity, and air temperature, as predictors of the foundation model residual.

This residual correction view appears empirically in regional SPEI forecasting. Figure 1(a) shows that the frozen TimesFM model produces conservative predictions and dampens extreme SPEI fluctuations during drought events. Figure 1(b) plots TimesFM residuals against the within event temperature anomaly, defined as the month to month change in maximum temperature in the same episode. In hotter than normal months, the model tends to underpredict drought severity, so part of its residual moves with the anomaly and is correctable. Table 4 reports the corresponding correlations between residuals and physical forcing.

Based on the example in Figure 1, we formulate test time adaptation of frozen foundation models for regional drought prediction as conditional residual modeling. If a TSFM has a regional sensitivity mismatch in its SPEI prediction under precipitation or temperature changes, then part of its one step residual is physically informed and predictable. Residual variations not explained by recent target history or physical forcing are treated as stochastic and should not be amplified. PhyTTA therefore corrects the frozen forecast only when the

current feature regime is close to recently observed regimes. It realizes this through three components: a *dynamic* forcing corrector, a *static* background bias tracker, and a *conditional* confidence gate.

In summary, our contributions are as follows:

- We formalize test time adaptation of frozen foundation models for regional drought prediction as physics guided conditional residual correction.
- We introduce the PhyTTA framework, which corrects the predictable part of the residual through three complementary components.
- We provide a formal analysis of the framework and show up to 19.78% MSE reduction in evaluations of four TSFM backbones on South Australian locations and the TimesFM backbone on cross regional tests.

2 Problem Setting

Let $y_t \in \mathbb{R}$ be the target SPEI value and $\mathbf{c}_t \in \mathbb{R}^d$ be the climate covariate vector at time t . We define the input $\mathbf{X}_t$ as the concatenation of historical targets and covariates over a lookback window L :

$$\mathbf{X}_t = [y_{t-L+1:t}, \mathbf{c}_{t-L+1:t}]. \quad (1)$$

We focus on one step ahead forecasting ($H = 1$) with a frozen TSFM f_θ :

$$\hat{y}_{t+1}^{(0)} = f_\theta(\mathbf{X}_t), \quad (2)$$

where θ denotes frozen backbone parameters and the superscript (0) denotes the base uncorrected forecast. Our goal is to produce a corrected forecast $\hat{y}_{t+1}$ without updating θ . The state vector S_t in Section 3.2 is a compact target history summary derived from $\mathbf{X}_t$ and used only by the residual corrector.

The quantity PhyTTA corrects is the one step residual $r_{t+1} = y_{t+1} - \hat{y}_{t+1}^{(0)}$, and $\mathbb{E}[\cdot]$ denotes expectation. At test time, y_{t+1} is never used before predicting $\hat{y}_{t+1}$. After the observation becomes available, it is used only to update the residual buffer for future forecasts. Thus, PhyTTA follows a rolling delayed feedback protocol, in which each label arrives one step after its forecast, rather than label free test time adaptation [Wang *et al.*, 2021; Kim *et al.*, 2025].

3 Method

3.1 The PhyTTA Framework

As shown in Figure 2, PhyTTA keeps the foundation model frozen and corrects its forecast along two paths. The top path maps the historical context to the base forecast $\hat{y}_{t+1}^{(0)}$. The bottom path forms two correction signals: a *dynamic* forcing residual $\hat{r}_{t+1}$ from a ridge corrector (Section 3.2) and a *static* slowly varying bias $\hat{B}_t$ from an exponential moving average tracker (Section 3.3). A *conditional* confidence gate keeps the correction only in supported regimes, and a bounded blend merges the two signals into the adjustment added to $\hat{y}_{t+1}^{(0)}$ (Section 3.4).

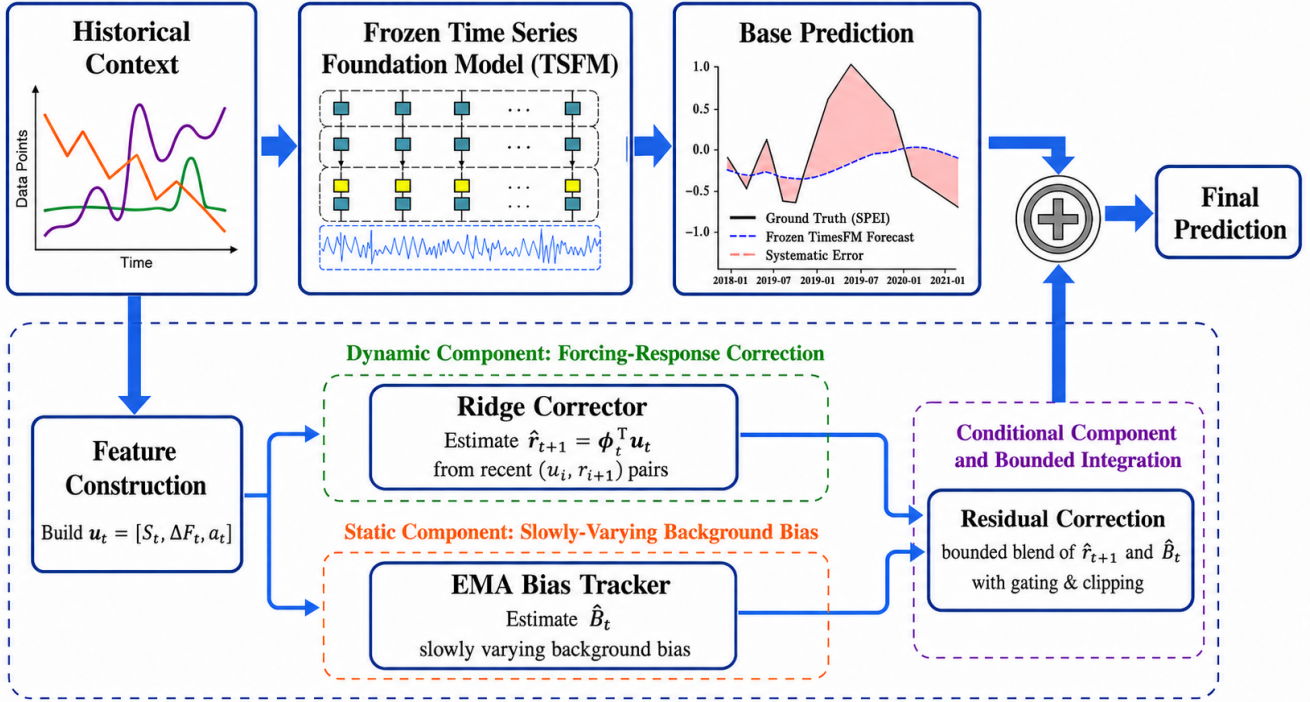


Figure 2: Overview of the PhyTTA framework. The frozen time series foundation model (TSFM) produces the base forecast $\hat{y}_{t+1}^{(0)}$. PhyTTA constructs $\mathbf{u}_t = [S_t, \Delta F_t, a_t]$ and estimates a dynamic forcing residual $\hat{r}_{t+1}$ with ridge regression and a static slowly varying bias $\hat{B}_t$ with an exponential moving average (EMA) tracker. The residual correction module combines them through a gated and bounded blend and adds the resulting correction to the base forecast to obtain $\hat{y}_{t+1}$.

3.2 Dynamic Component: Forcing Correction

The dynamic component in Figure 2 corrects the part of the residual correlated with recent forcing changes. It is a single linear ridge corrector over

$$\mathbf{u}_t = [S_t, \Delta F_t, a_t], \quad a_t = y_t, \quad (3)$$

built from the drought state S_t , current forcing change ΔF_t , and latest value $a_t = y_t$. The state and forcing blocks play complementary roles:

- The *system state* (S_t) captures the system’s internal inertia and “memory.” Since drought is an autoregressive process, the magnitude of the prediction error is often state dependent.
- The *physical forcing* (ΔF_t) represents the external drivers (precipitation and temperature changes) that drive the system’s evolution.

System State Features (S_t). We capture drought inertia using target history statistics. S_t contains full context descriptors over L and short window descriptors over $w = 16$ months, a short window spanning slightly more than one annual cycle:

$$S_t = [\mu_t^L, \sigma_t^L, y_t - y_{t-L+1}, \mu_{\Delta y, t}^L, \sigma_{\Delta y, t}^L, \mu_t^w, \sigma_t^w, y_t - y_{t-w+1}], \quad w = 16, \quad (4)$$

where μ and σ are the rolling mean and standard deviation over the window, applied to the target y and to its one step change $\Delta y_i = y_i - y_{i-1}$.

Physical Forcing Features (ΔF_t). From $\mathbf{c}_t$, we extract precipitation and temperature, the two primary physical drivers of SPEI, and two auxiliary atmospheric covariates used in the implementation:

$$\Delta F_t = [\Delta P_t, \Delta T_t^{\max}, \Delta Q_t, \Delta T_t^{\text{air}}], \quad (5)$$

where P_t denotes monthly precipitation, $\Delta P_t = P_t - P_{t-1}$, $\Delta T_t^{\max} = T_t^{\max} - T_{t-1}^{\max}$, $\Delta Q_t = Q_t - Q_{t-1}$, and $\Delta T_t^{\text{air}} = T_t^{\text{air}} - T_{t-1}^{\text{air}}$. Here $T_t^{\max}$ denotes the monthly value derived from daily maximum near surface temperature, Q_t denotes monthly specific humidity, and T_t^{air} denotes monthly mean near surface air temperature, all taken from NCEP-NCAR Reanalysis 1 [Kalnay *et al.*, 1996]. Because ΔF_t is a one step difference, it lets the corrector respond to immediate forcing changes.

Ridge corrector. The dynamic correction is the ridge prediction $\hat{r}_{t+1} = \phi_t^\top \mathbf{u}_t$, where ϕ_t is the ridge weight vector. PhyTTA keeps a buffer of the most recent N feature residual pairs,

$$\mathcal{W}_t = \{(\mathbf{u}_i, r_{i+1})\}_{i=t-N}^{t-1}, \quad (6)$$

updated only after each forecast is made, using $r_{t+1} = y_{t+1} - \hat{y}_{t+1}^{(0)}$. We fit ϕ_t by ridge regression on the pairs $(\mathbf{u}_i, r_{i+1})$ with penalty $\gamma \|\phi\|_2^2$, which has a closed form solution. We

use $N = 100$, warm up $N_{\min} = 30$, and $\gamma = 0.5$; before warm up, PhyTTA returns the base forecast. The short buffer lets ϕ_t update with recent conditions, and Eq. (9) combines this fast estimate with the slow background estimate.

3.3 Static Component: Slowly Varying Background Bias

The static component, the lower branch in Figure 2, tracks the slow bias signal $\hat{B}_t$. The dynamic corrector targets residual variations that move with recent state and forcing changes, but it cannot reliably remove a persistent offset between the frozen forecast and the observed target. Let B_t denote the latent low frequency component of the frozen model residual, representing a slowly varying regional offset between the base forecast and the observed target. Since B_t is not directly observed, we estimate its running value from past residuals using a separate slow channel, an exponential moving average (EMA):

$$\hat{B}_{t+1} = \alpha_B r_{t+1} + (1 - \alpha_B) \hat{B}_t. \quad (7)$$

Here $\hat{B}_t$ is the exponential moving average estimate of the latent low frequency bias B_t , and $\alpha_B = 0.3$.

The EMA is a low pass filter: it smooths fast residual fluctuations and tracks a slowly moving baseline. Since each update retains a $(1 - \alpha_B)$ share of the previous estimate, isolated residual outliers have limited influence. Thus, “static” means slow relative to the fast ridge corrector, not time invariant. This slow offset differs from two other slow quantities in the method: the state memory inside S_t , which is only an input to the dynamic corrector, and the temporal correlation of the forcing changes themselves.

3.4 Conditional Component and Bounded Integration

The dynamic and static estimates are candidate correction signals; neither is always safe to apply. PhyTTA therefore uses a support gate for unfamiliar regimes and a bounded blend to cap accepted corrections.

Confidence gate. We measure regime support by the nearest neighbor distance to the residual buffer in standardized feature space:

$$q_t = \frac{1}{1 + \min_{\mathbf{u}_i \in \mathcal{W}_t} \|\mathbf{u}_t - \mathbf{u}_i\|_2}. \quad (8)$$

Here $q_t \in (0, 1]$ is a support score, with larger values meaning closer buffered support. The score is used only after $|\mathcal{W}_t| \geq N_{\min}$. If $q_t > \tau$, PhyTTA applies the residual correction; otherwise it returns the frozen forecast $\hat{y}_{t+1}^{(0)}$. We set $\tau = 0.10$ after checking $\{0.05, 0.10, 0.15\}$ on validation; no test set tuning is used. Feature z scores are fixed from the training and warm up portion.

Bounded blend. When the main gate is open, the final forecast combines $\hat{r}_{t+1}$ and $\hat{B}_t$, then clips the result. The static estimate is included only when the stricter bias gate $q_t > \tau_B$ is satisfied. When the main gate is closed, the correction is

zero:

$$\begin{aligned} \omega_t &= \omega \mathbb{I}(q_t > \tau_B), \\ \bar{g}_t &= \text{clip}\left((1 - \omega_t)\hat{r}_{t+1} + \omega_t\hat{B}_t, -c_r, c_r\right), \\ \hat{y}_{t+1} &= \begin{cases} \hat{y}_{t+1}^{(0)}, & q_t \leq \tau, \\ \text{clip}\left(\hat{y}_{t+1}^{(0)} + \lambda\bar{g}_t, -c_y, c_y\right), & q_t > \tau. \end{cases} \end{aligned} \quad (9)$$

Here $\text{clip}(x, a, b) = \min\{\max\{x, a\}, b\}$ and $\mathbb{I}(\cdot)$ equals 1 when its condition is true and 0 otherwise. Constants are held fixed across backbones and regions. The step size $\lambda = 0.8 < 1$ is conservative, and $\omega = 0.3 < 0.5$ keeps the fast forcing term dominant over the slowly varying bias. The clips are safety rails on the standardized SPEI scale: $c_y = 3.5$ avoids truncating plausible forecasts, while $c_r = 1.5$ caps any single residual correction. The bias gate $\tau_B = 0.60 > \tau$ controls only whether the EMA bias term enters the blend.

Online update order. The delayed feedback protocol in Section 2 is implemented as follows. At each step t , PhyTTA first computes the frozen forecast $\hat{y}_{t+1}^{(0)} = f_\theta(X_t)$ from the available history and constructs $\mathbf{u}_t$. If $|\mathcal{W}_t| < N_{\min}$, it returns the frozen forecast. Otherwise, it fits the ridge corrector on $\mathcal{W}_t$, computes $\hat{r}_{t+1}$, q_t , $\bar{g}_t$, and the corrected forecast $\hat{y}_{t+1}$ by Eq. (9). In both cases, y_{t+1} is used only after it becomes available: the method then computes $r_{t+1} = y_{t+1} - \hat{y}_{t+1}^{(0)}$, updates $\mathcal{W}_{t+1}$, and updates the EMA estimate $\hat{B}_{t+1}$. Thus, each newly observed label affects only later forecasts.

4 Theoretical Analysis

Our analysis is conditional on the existence of a predictable component in the frozen model residual. It does not claim that every finite buffer online update improves every test point. Instead, it shows that if the residual contains a component predictable from the features already used by PhyTTA, then correcting this component reduces expected squared error.

4.1 Why Residual Correction Helps

Recall that $r_{t+1} = y_{t+1} - \hat{y}_{t+1}^{(0)}$ is the one step residual of the frozen backbone and $\mathbf{u}_t = [S_t, \Delta F_t, a_t]$ is the feature vector used by the residual corrector. We first analyze the population target estimated by the online ridge corrector.

Proposition 1 (Predictable residual correction reduces error). *Let $\hat{r}_{t+1} = \phi_t^\top \mathbf{u}_t$ be the best penalized linear fit of r_{t+1} from $\mathbf{u}_t$, corresponding to the population version of the ridge corrector in Section 3.2. Then*

$$\mathbb{E}[(r_{t+1} - \hat{r}_{t+1})^2] \leq \mathbb{E}[r_{t+1}^2]. \quad (10)$$

The inequality is strict whenever r_{t+1} contains a nonzero feature correlated component that is not fully shrunk to zero by the ridge penalty.

Proof. The zero correction is always feasible: setting $\phi_t = 0$ gives $\hat{r}_{t+1} = 0$, which recovers the frozen forecast. Since $\hat{r}_{t+1}$ is the best penalized linear fit, its penalized error cannot exceed that of zero correction:

$$\mathbb{E}[(r_{t+1} - \hat{r}_{t+1})^2] + \gamma \|\phi_t\|_2^2 \leq \mathbb{E}[r_{t+1}^2]. \quad (11)$$

Because the ridge penalty is nonnegative, the unpenalized squared error after correction is also no larger than $\mathbb{E}[r_{t+1}^2]$. Strict reduction occurs when a nonzero predictable residual component is fitted rather than suppressed. $\square$

This result explains the target of the dynamic ridge component, not an unconditional guarantee for every finite buffer estimate. In practice, the estimate can be noisy when the current feature regime is poorly supported by recent history. This motivates the confidence gate, shrinkage factor $\lambda < 1$, and clipping in Eq. (9): the correction is applied only when $q_t > \tau$ and its magnitude is bounded.

The EMA term targets a different predictable component: a slowly varying residual offset. It is not treated as the intercept of the ridge proof. Instead, Eq. (7) estimates a low frequency bias $\hat{B}_t$ from past residuals only. If recent residuals share a persistent sign, $\hat{B}_t$ is aligned with the next residual and can reduce squared error by the same principle. The stricter gate $\tau_B > \tau$ reflects that this bias term shifts the forecast baseline and should be used only in strongly supported regimes.

4.2 When the Premise Holds for SPEI

The above argument is useful only when the frozen model residual is predictable from $\mathbf{u}_t$. This premise is natural in regional SPEI forecasting.

Corollary 1 (SPEI residual predictability). *Suppose that the one step residual after subtracting the frozen backbone forecast, $r_{t+1} = y_{t+1} - \hat{y}_{t+1}^{(0)}$, contains a nonzero component that is linearly predictable from the forecast time features $\mathbf{u}_t = [S_t, \Delta F_t, a_t]$. In regional SPEI forecasting, such a component can arise from a forcing mismatch with respect to precipitation change, temperature change, or recent drought state. Then the population residual correction in Proposition 1 reduces the expected squared error, with strict reduction when this predictable component is not fully shrunk to zero by the ridge penalty.*

SPEI reflects the climatic water balance: precipitation contributes to water supply, while temperature affects evaporative demand [Vicente-Serrano *et al.*, 2010; Beguería *et al.*, 2014]. A frozen foundation model trained on heterogeneous time series may learn a useful average forecasting rule but still misrepresent local forcing sensitivity. When this happens, the missed regional response remains in the one step residual and can move with ΔF_t , S_t , or the anchor $a_t = y_t$.

The sign of this association between residuals and forcing is not interpreted as the physical sign of the SPEI response itself. It reflects whether the frozen model underestimates or overestimates the local forcing response. Our claim is predictive: if the residual is associated with observable forecast time features, then a residual corrector can use that association.

Monthly drought dynamics further support this premise because SPEI has persistence through precipitation deficits, evaporative demand, soil moisture memory, and large scale climate modes [Dai, 2013; Trenberth *et al.*, 2014]. Therefore, recent drought state and current forcing changes contain information about the next residual. The EMA tracker handles the complementary case of slow background bias, where

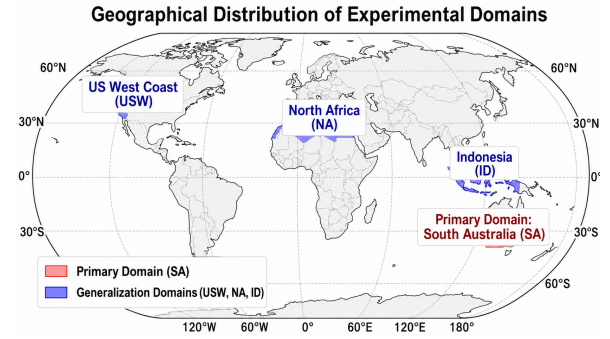


Figure 3: Geographic distribution of the four climate regions. Primary region: South Australia (SA), bounded from 26.0°S to 38.0°S and from 129.0°E to 141.0°E. Generalization regions: US West Coast (USW), North Africa (NA), and Indonesia (ID).

residuals remain persistently positive or negative over several months.

Table 4 empirically checks this residual predictability premise by reporting correlations between residuals and forcing across regions. Table 3 further shows that physical forcing is important: removing ΔF_t turns the method into unguided residual fitting and can degrade performance below the frozen baseline.

5 Experiments

5.1 Study Area

As shown in Figure 3, we evaluate PhyTTA using one primary region and three generalization regions.

Primary Region: South Australia (SA). We choose South Australia as the primary region due to its hydroclimatic variability and sensitivity to large scale modes such as the El Niño-Southern Oscillation (ENSO) and the Indian Ocean Dipole (IOD); we focus on its agricultural zone (bounds in Figure 3).

Generalization Regions. We further evaluate on three regions: (i) **US West Coast (USW)**, a Mediterranean regime where temperature anomalies can intensify drought through elevated evaporative demand [Dai, 2013]; (ii) **North Africa (NA)**, an arid region with strong low frequency regional offsets and limited effective data; and (iii) **Indonesia (ID)**, a tropical monsoon regime dominated by precipitation driven drought dynamics.

5.2 Datasets

We evaluate monthly prediction of the 30-day SPEI series from SPEI-GD [Liu *et al.*, 2024a]. The study period spans 1982/01 to 2018/12. The dataset contains 39 features per time step:

- **Target:** 30-day SPEI from SPEI-GD [Liu *et al.*, 2024a], sampled monthly.
- **Meteorological covariates:** 34 variables from the National Centers for Environmental Prediction-National Center for Atmospheric Research (NCEP-NCAR) Reanalysis 1 [Kalnay *et al.*, 1996], including atmospheric dynamics and surface conditions.

Method	Loc. 1			Loc. 2			Loc. 3		
	MSE ↓	MAE ↓	R^2 ↑	MSE ↓	MAE ↓	R^2 ↑	MSE ↓	MAE ↓	R^2 ↑
Classical and Recurrent Baselines									
ARIMA	1.213	0.882	-0.417	1.201	0.875	-0.409	1.225	0.887	-0.426
ETS	0.950	0.765	0.050	0.942	0.760	0.055	0.957	0.769	0.048
LSTM	0.688	0.612	0.320	0.679	0.605	0.326	0.694	0.616	0.318
GRU	0.674	0.598	0.335	0.667	0.592	0.340	0.681	0.603	0.330
TCN	0.662	0.586	0.345	0.655	0.581	0.350	0.670	0.590	0.340
Deep Learning SOTA									
Transformer	0.537	0.554	0.451	0.510	0.581	0.522	0.745	0.655	0.365
Autoformer	0.546	0.559	0.442	0.547	0.582	0.487	0.752	0.658	0.358
Crossformer	0.542	0.557	0.446	0.474	0.588	0.556	0.740	0.654	0.370
DLinear	0.533	0.551	0.455	0.423	0.553	0.604	0.735	0.650	0.375
FiLM	0.558	0.569	0.430	0.529	0.597	0.505	0.760	0.662	0.350
iTransformer	0.524	0.546	0.464	0.454	0.539	0.575	0.738	0.652	0.372
PatchTST	0.535	0.552	0.453	0.461	0.548	0.568	0.743	0.655	0.366
TimesNet	0.550	0.562	0.438	0.489	0.572	0.542	0.754	0.659	0.356
TSMixer	0.529	0.549	0.459	0.468	0.555	0.561	0.736	0.651	0.374
Foundation Models with Test-Time Adaptation									
TimesFM v2.5 (Base)	0.391	0.466	0.564	0.288	0.385	0.734	0.365	0.448	0.611
TimesFM v2.5 + RevIN	0.373	0.453	0.584	0.279	0.377	0.742	0.345	0.430	0.633
TimesFM v2.5 + TAFAS	0.350	0.434	0.610	0.265	0.367	0.756	0.325	0.414	0.654
TimesFM v2.5 + PhyTTA (Ours)	0.343	0.435	0.618	0.249	0.362	0.770	0.310	0.416	0.670
Chronos-1 (Base)	0.398	0.456	0.558	0.304	0.379	0.718	0.442	0.466	0.532
Chronos-1 + RevIN	0.415	0.459	0.540	0.296	0.374	0.726	0.433	0.458	0.540
Chronos-1 + TAFAS	0.406	0.463	0.549	0.298	0.381	0.724	0.413	0.459	0.560
Chronos-1 + PhyTTA (Ours)	0.362	0.446	0.598	0.245	0.356	0.772	0.360	0.418	0.622
TabPFN-ts v1.0.8 (Base)	0.472	0.503	0.532	0.444	0.490	0.550	0.411	0.471	0.573
TabPFN-ts v1.0.8 + RevIN	0.382	0.448	0.574	0.290	0.358	0.733	0.361	0.430	0.615
TabPFN-ts v1.0.8 + TAFAS	0.372	0.441	0.585	0.286	0.372	0.736	0.353	0.429	0.624
TabPFN-ts v1.0.8 + PhyTTA (Ours)	0.361	0.445	0.597	0.238	0.334	0.781	0.373	0.430	0.600
Moirai (Base)	0.402	0.469	0.552	0.309	0.386	0.712	0.462	0.476	0.508
Moirai + RevIN	0.522	0.568	0.418	0.672	0.549	0.382	0.494	0.559	0.474
Moirai + TAFAS	0.454	0.531	0.494	0.398	0.475	0.632	0.437	0.518	0.534
Moirai + PhyTTA (Ours)	0.382	0.456	0.572	0.232	0.330	0.785	0.360	0.416	0.620

Table 1: One month ahead SPEI forecasting performance across three South Australian locations: Loc. 1 (-26.125° , 129.125° E), Loc. 2 (-29.125° , 134.875° E), and Loc. 3 (-35.625° , 138.875° E). Lower MSE/MAE is better and higher R^2 is better. Bold values denote the best value within each foundation backbone block.

- **Climate indices:** the Southern Oscillation Index (SOI) [Trenberth, 1984], the Indian Ocean Dipole (IOD) [Saji *et al.*, 1999], the Southern Annular Mode (SAM) [Marshall, 2003], and the Niño 3.4 ENSO index [Trenberth, 1997].

We split each series chronologically: the first 70% for training and historical warm up, the next 10% for validation, and the final 20% for testing. The confidence gate τ is selected on the validation split (Sections 3.2 to 3.4) with all remaining scalar settings kept fixed, and reported metrics are computed on the held out test split.

5.3 Baselines and Implementation Sources

To ensure a controlled comparison, we follow the one step rolling protocol of TAFAS [Kim *et al.*, 2025] and compare against classical, neural, foundation model, and test time adaptation baselines. The implementation sources are fixed as follows: ARIMA and ETS use `statsmodels` [Seabold and

Perktold, 2010]; recurrent and neural baselines use PyTorch [Paszke *et al.*, 2019] and the public Time Series Library style implementations [Wu *et al.*, 2023a]; foundation model baselines use the public releases of TimesFM v2.5 [Google Research, 2025], Chronos, Moirai, and TabPFN with the tabPFN time series adapter v1.0.8 [Prior Labs, 2025]. The released repository¹ contains the PhyTTA correction module and the TimesFM runner; auxiliary scripts for the additional foundation backbones, the validation/test split protocol of Section 5.2, and per table reproduction commands are included in the code repo. The following list specifies the baseline families used in Table 1.

1. Classical and recurrent models: ARIMA [Box *et al.*, 2015], ETS [Hyndman *et al.*, 2002], LSTM [Hochreiter and Schmidhuber, 1997], GRU [Cho *et al.*, 2014], and TCN [Bai *et al.*, 2018].

2. Deep learning SOTA: Transformer [Vaswani *et al.*, 2017],

¹<https://github.com/Wentao-Gao/PhyTTA>

Autoformer [Wu *et al.*, 2021], Crossformer [Zhang and Yan, 2023], DLinear [Zeng *et al.*, 2023], FiLM [Zhou *et al.*, 2022], iTransformer [Liu *et al.*, 2024b], PatchTST [Nie *et al.*, 2023], TimesNet [Wu *et al.*, 2023b], and TSMixer [Chen *et al.*, 2023].

3. Latest Foundation models: TabPFN [Hollmann *et al.*, 2025] with the `tabpfn_time_series` adapter (tabpfn-ts v1.0.8) [Prior Labs, 2025], Chronos [Ansari *et al.*, 2024], Moirai [Woo *et al.*, 2024], and TimesFM v2.5 [Das *et al.*, 2024; Google Research, 2025].

4. Test time adaptation: RevIN [Kim *et al.*, 2022] and TAFAS [Kim *et al.*, 2025].

Experimental Protocol. We use rolling one step ahead forecasting ($H = 1$) with context length $L = 288$ months and residual buffer $N = 100$; backbone parameters remain frozen, with hyperparameters as in Eq. (9) and Section 3.2. Within block comparisons compare each frozen base backbone with its PhyTTA corrected version on a shared window; cross block numbers use different windows and are shown for orientation only.

5.4 Performance Evaluation

Main results. Table 1 reports one step ahead ($H = 1$) SPEI forecasting on three representative SA locations (coordinates in the caption).

PhyTTA achieves 12.3% to 15.1% MSE reductions for TimesFM across the three SA locations with consistent R^2 gains, and improves the frozen Chronos, TabPFN, and Moirai backbones at every location. Together with TimesFM, these are 12 backbone location comparisons. RevIN gives modest gains, and TAFAS, which needs inference time gradient updates, edges PhyTTA on MAE at TimesFM Locations 1 and 3 (and beats it on TabPFN Loc. 3). Moirai+RevIN degrades sharply while Moirai+PhyTTA recovers, indicating that physics guided correction is more robust than generic normalization under regional shift.

Cross Regional Generalization. Table 2 shows the evaluation results of TimesFM + PhyTTA on three regions with distinct drought mechanisms. PhyTTA improves MSE in all three, with the largest gains on US West Coast and North Africa and a smaller positive gain on precipitation dominated Indonesia, supporting our hypothesis that correctable residual magnitude depends on regional physical characteristics rather than architecture; relative MSE reduction is the more comparable cross region quantity given different SPEI variance across regions.

Region	Base (TimesFM)			PhyTTA			Imp.
	MSE ↓	MAE ↓	R^2 ↑	MSE ↓	MAE ↓	R^2 ↑	
US West (USW)	0.0546	0.1770	0.9407	0.0438	0.1565	0.9526	+19.78%
North Africa (NA)	0.0695	0.1953	0.9291	0.0590	0.1839	0.9398	+15.11%
Indonesia (ID)	0.2302	0.3827	0.6207	0.2212	0.3705	0.6354	+3.91%

Table 2: Cross regional generalization of PhyTTA with TimesFM backbone. Imp. denotes relative MSE reduction (Base to PhyTTA).

5.5 Analysis

Ablation Study. Table 3 isolates the physics and bias components on SA Location 1. Full PhyTTA performs best;

Variant	MSE ↓	Improvement ↑	Analysis
Base Model	0.3910		Baseline
w/o Physics	0.4080	-4.35%	Statistical residual fitting only
w/o Bias	0.3597	+8.01%	No EMA component
Full PhyTTA	0.3428	+12.33%	Best on this location

Table 3: Ablation study on SA Location 1. The w/o Physics variant removes ΔF_t while retaining the online residual learner and state features; the result worse than the frozen baseline indicates that unguided residual fitting can overfit noise.

Region	Corr($r_{t+1}, \Delta P_t$)	Corr($r_{t+1}, \Delta T_t^{\max}$)	Observed Sensitivity
South Australia	-0.31	+0.28	Rainfall sensitive
US West Coast	-0.23	+0.09	Moderate rainfall sensitivity
Indonesia	-0.43	-0.02	Precipitation dominated
North Africa	-0.41	+0.31	Mixed rain and temperature

Table 4: Pearson correlations between TimesFM residuals and physical forcing changes on the test set; SA row pools the three locations in Table 1, each generalization row is computed on its corresponding regional test set. Bold: $|\rho| \geq 0.3$.

removing physical forcing makes the corrected MSE worse than the frozen base, so naive residual fitting can harm a strong frozen model without physics guided inputs. Removing the EMA bias tracker leaves only a partial gain, confirming that the dynamic and static components are complementary.

Mechanism Verification. To support Corollary 1’s residual predictability premise, we correlate TimesFM residuals with forcing changes (Table 4). Residuals associate meaningfully with ΔP_t or $\Delta T_t^{\max}$ across regions: Indonesia is strongly precipitation driven, North Africa shows a mixed signature on both drivers, and South Australia also shows non trivial correlations on both, consistent with compound drought triggers.

6 Conclusion

We have presented PhyTTA, an inference time residual correction framework for frozen TSFMs on regional drought prediction, combining a dynamic forcing corrector, a static EMA background bias tracker, and a conditional confidence gate. With TimesFM, PhyTTA cuts MSE by 19.78%/15.11% on USW/NA and by 12.3% to 15.1% across three SA locations, and improves the frozen base in all 12 SA backbone location cells (best within block in 11/12). Future work will extend PhyTTA to multi step horizons and uncertainty aware warnings.

7 Social Impact

This work is informed by technical exchange with multidisciplinary teams at CSIRO Environment and CSIRO Technology. PhyTTA targets regional drought early warning settings, where more accurate one step ahead SPEI forecasts can support planning and risk management. Its design uses no backbone parameter updates for low latency deployment and physics guided features for interpretability. The forecasts should not be used as sole decision triggers; deployment requires uncertainty communication, expert review, and decision aligned evaluation.

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